

# Holocene variability of Benguela upwelling

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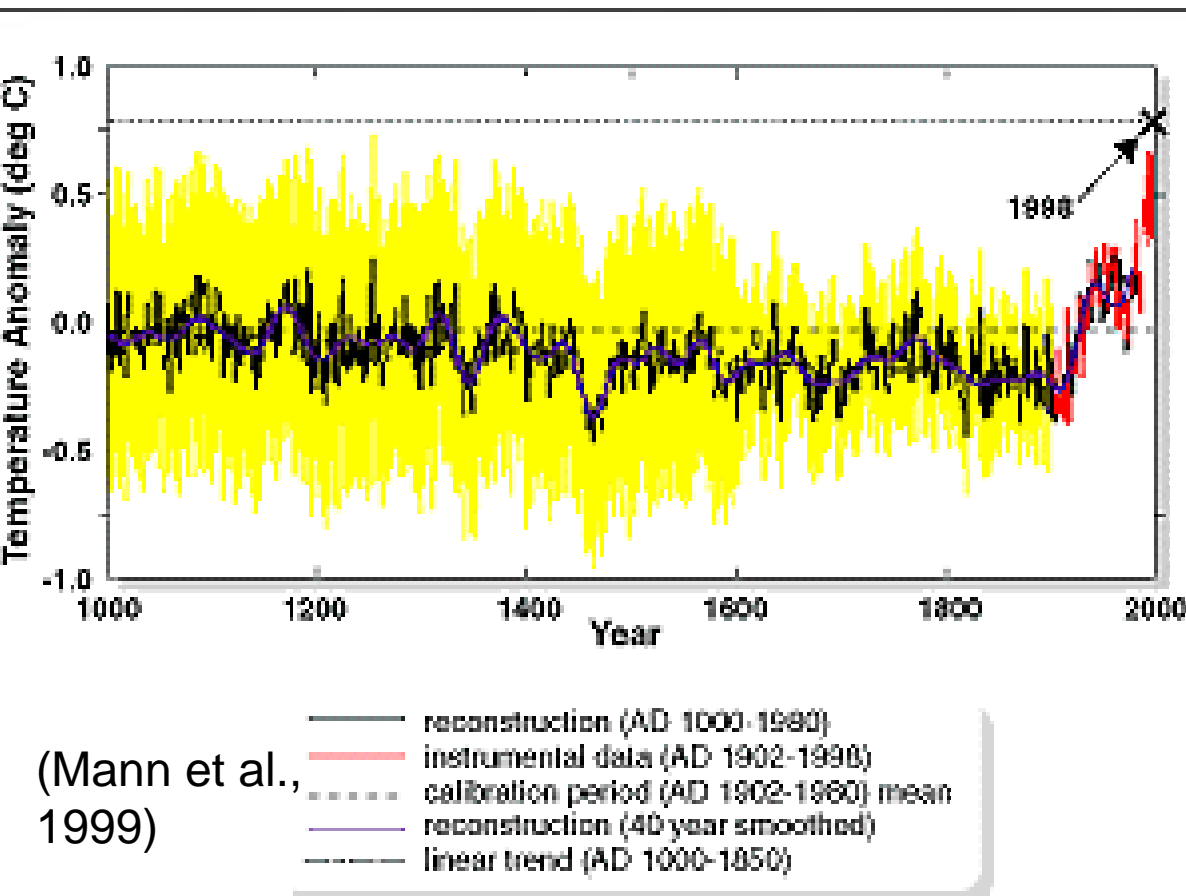
With thanks to: DOE/ORAU/GREF, LDEO Climate Center, ODP, Jean Lynch-Stieglitz, Martin Visbeck, Tom Guilderson, Tom Koutavas, Sarah Ingram, Martha Bryan, Linda Baker, Pat Malone, Grace Kim, and so many other helpful folks...



ODP's *JOIDES Resolution*

<http://www-odp.tamu.edu/>

# Why do we care about millennial climate change?



To understand historical climate change (red), we need to understand “natural” geological climate change (yellow)

# tropical Atlantic climate change over the last 20,000 years

- Greenland ice cores show little Holocene climate variability
- global marine records suggest Little Ice Age is part of a persistent pattern:
  - ★ how widespread is it ?
  - ★ origin in tropics? high-latitudes?
  - ★ S Hem same timing as N Hem?

(Dansgaard, 1993)

(Bond, 2001; deMenocal 2000)

# timing of deglacial climate change between hemispheres:

N Hem deglaciation  
cold interval

“Younger Dryas”  
(11.5-13 kyr BP)

lags comparable  
event in S Hem

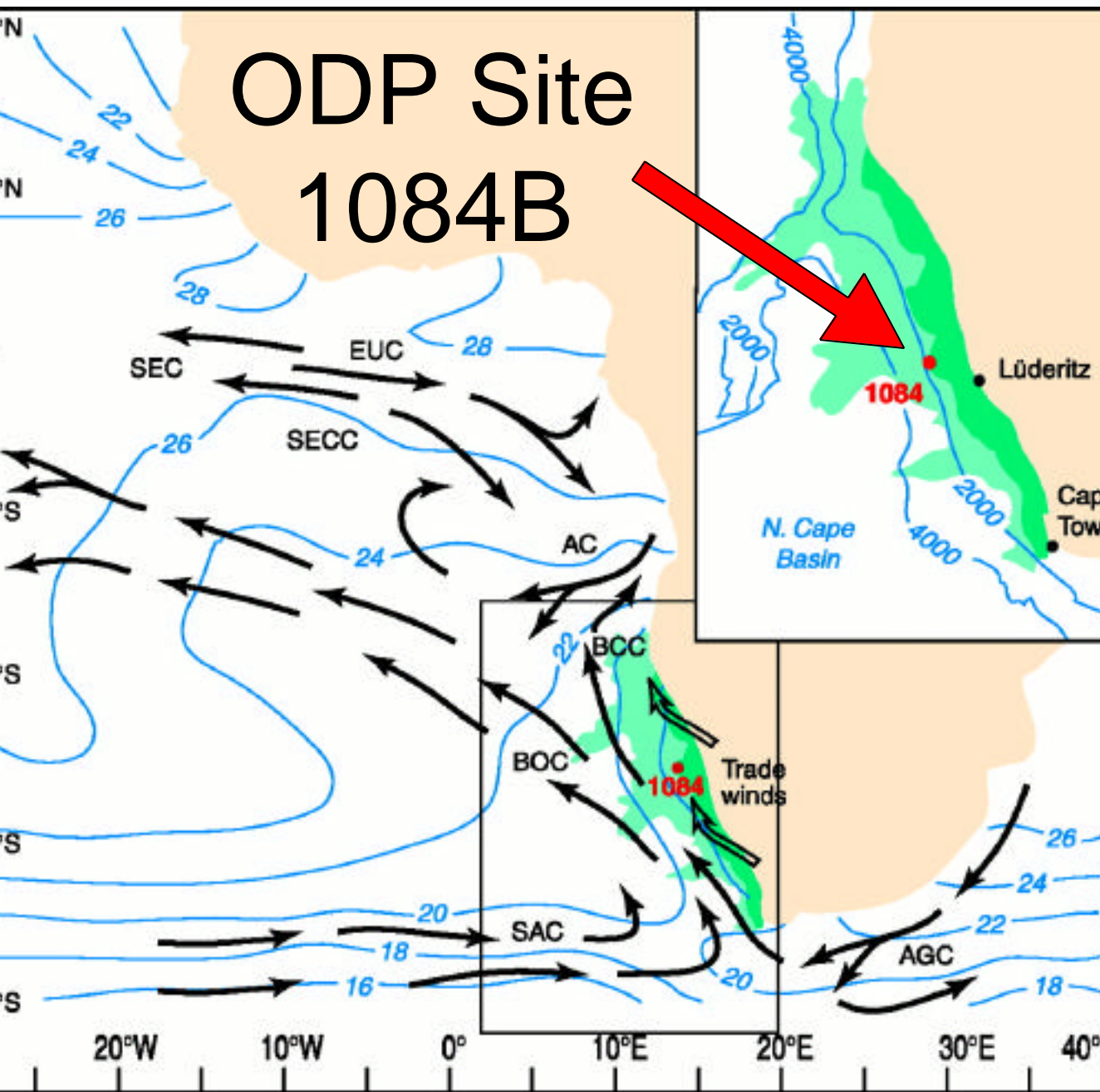
Blunier et al., 1998)

- N Hem deglacial lag  
interval confirmed:

~ 1500 yrs

(Charles et al., 1996)

- ★ S Hem subtropics  
match *N Hem* polar  
regions in deglacial  
timing!



Benguela  
upwelling  
zone off  
Namibia

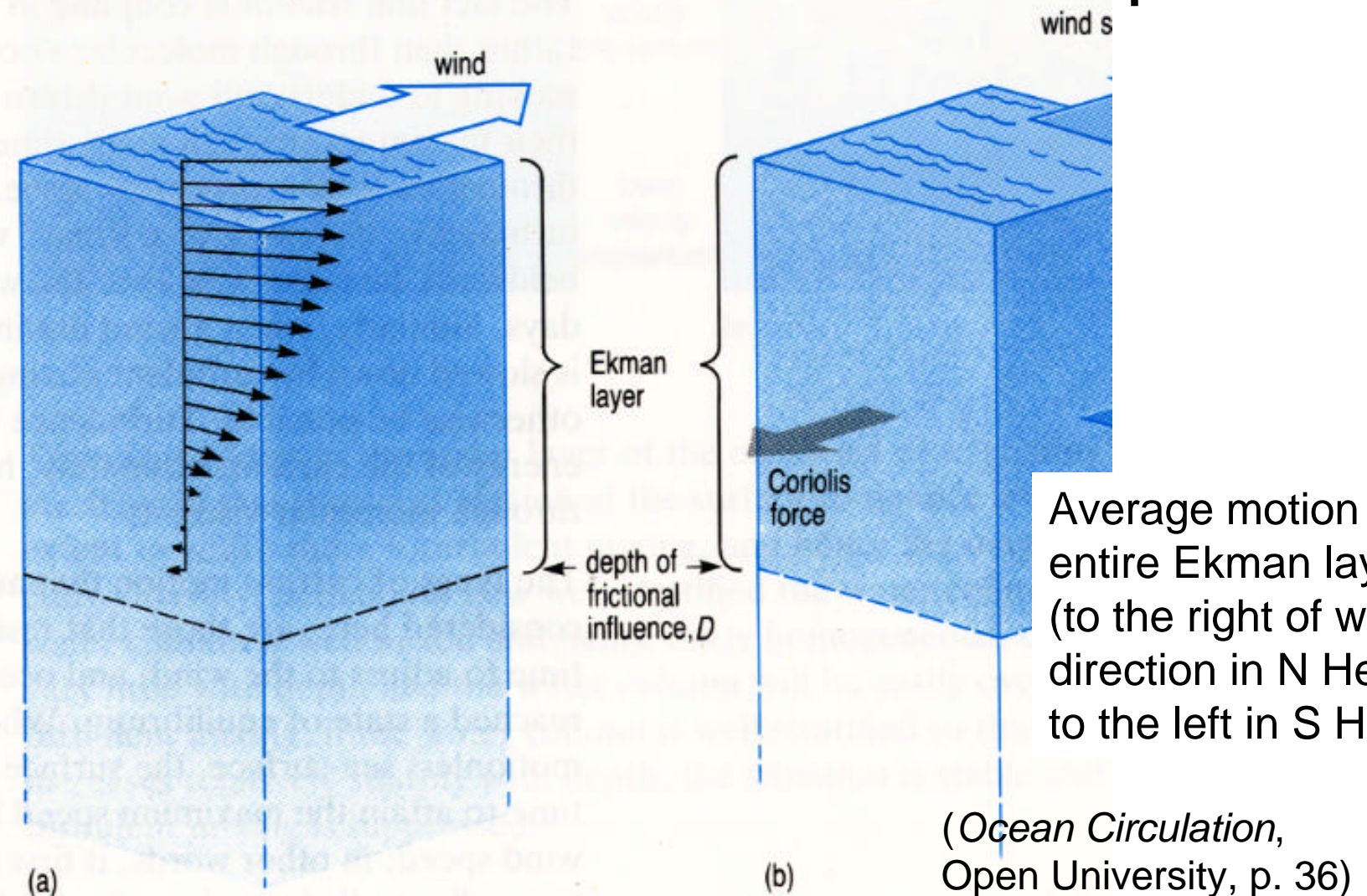
near where  
Indian Ocean  
water mixes  
with waters  
from South  
Atlantic

2000m water  
depth on  
continental  
shelf

(Marlow et al.  
2000, Science)



# physics of upwelling (in 5 min or less): wind stress induces 'Ekman spiral'



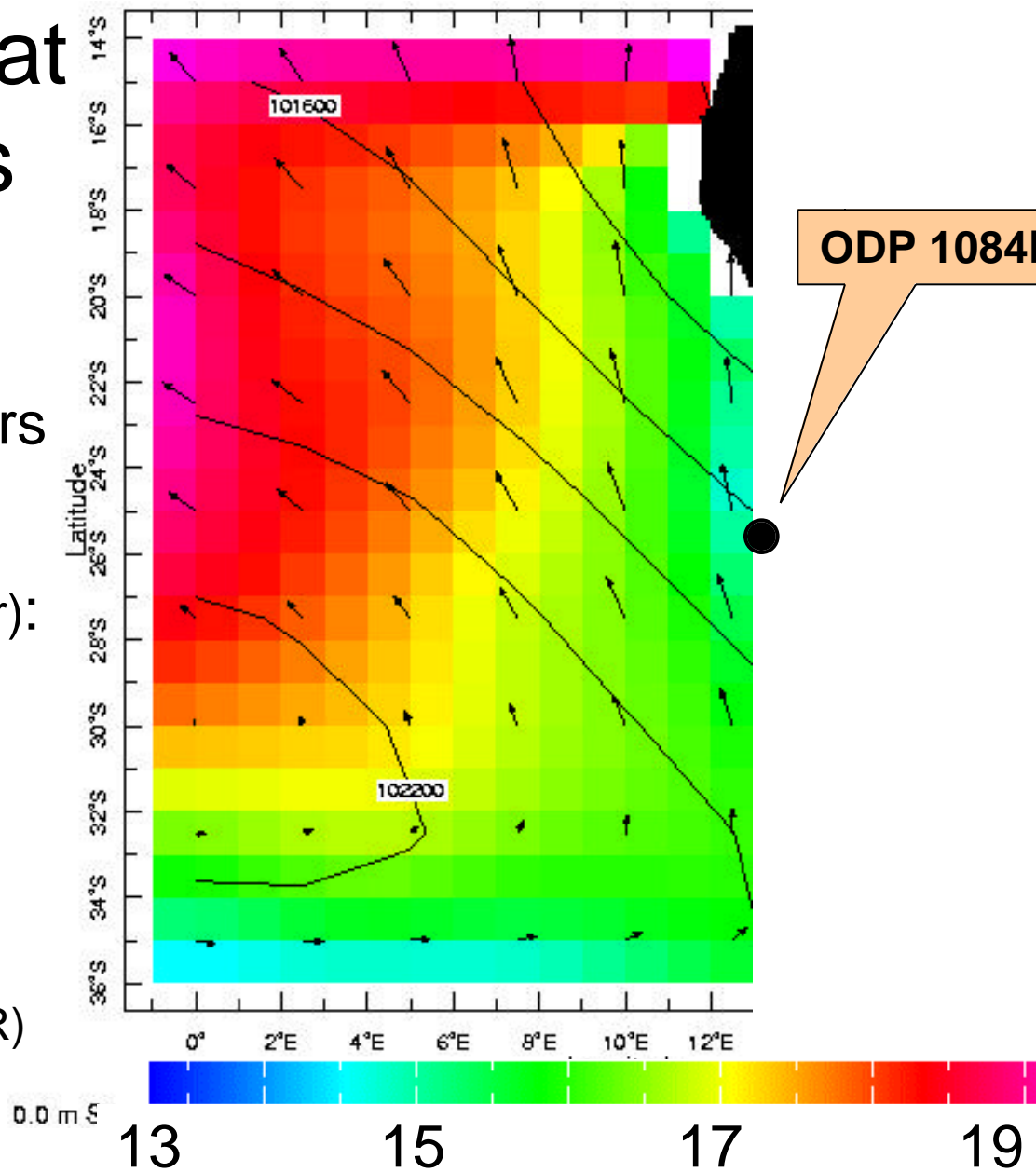
# Divergence at coast gives upwelling:

deep, colder waters  
rise to surface  
along coast

Feb (S Hem summer):  
least upwelling,  
SST=17.2 °C

Sep (S Hem winter):  
most upwelling,  
SST=14.8 °C

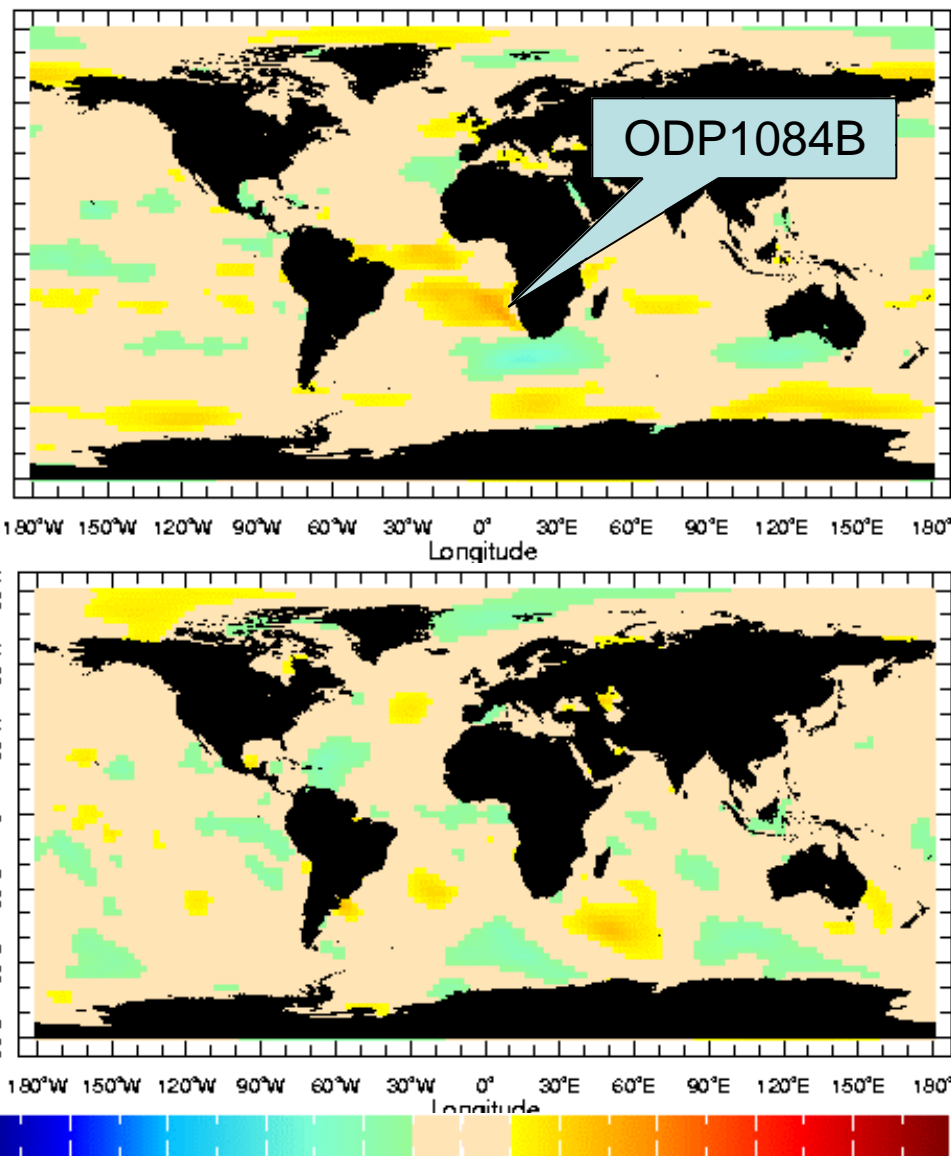
DAA WOA, NCEP-NCAR)



# correlations with global wind strength:

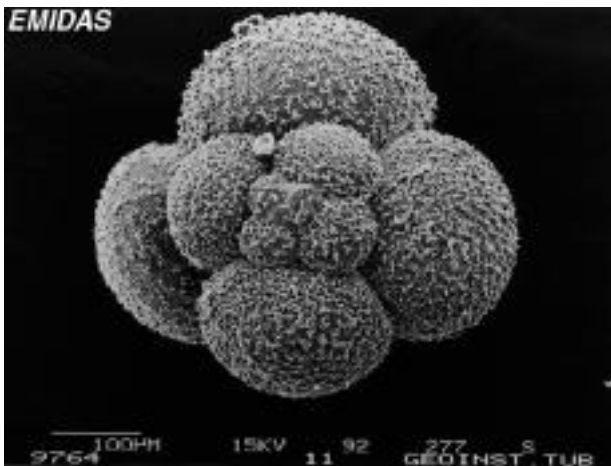
- monthly anomaly SST at site ODP1084B
- highest for local zonal wind strength  
15-20 % of variance

(NOAA WOA, NCEP-NCAR)





# *G. bulloides*: planktic foraminifera



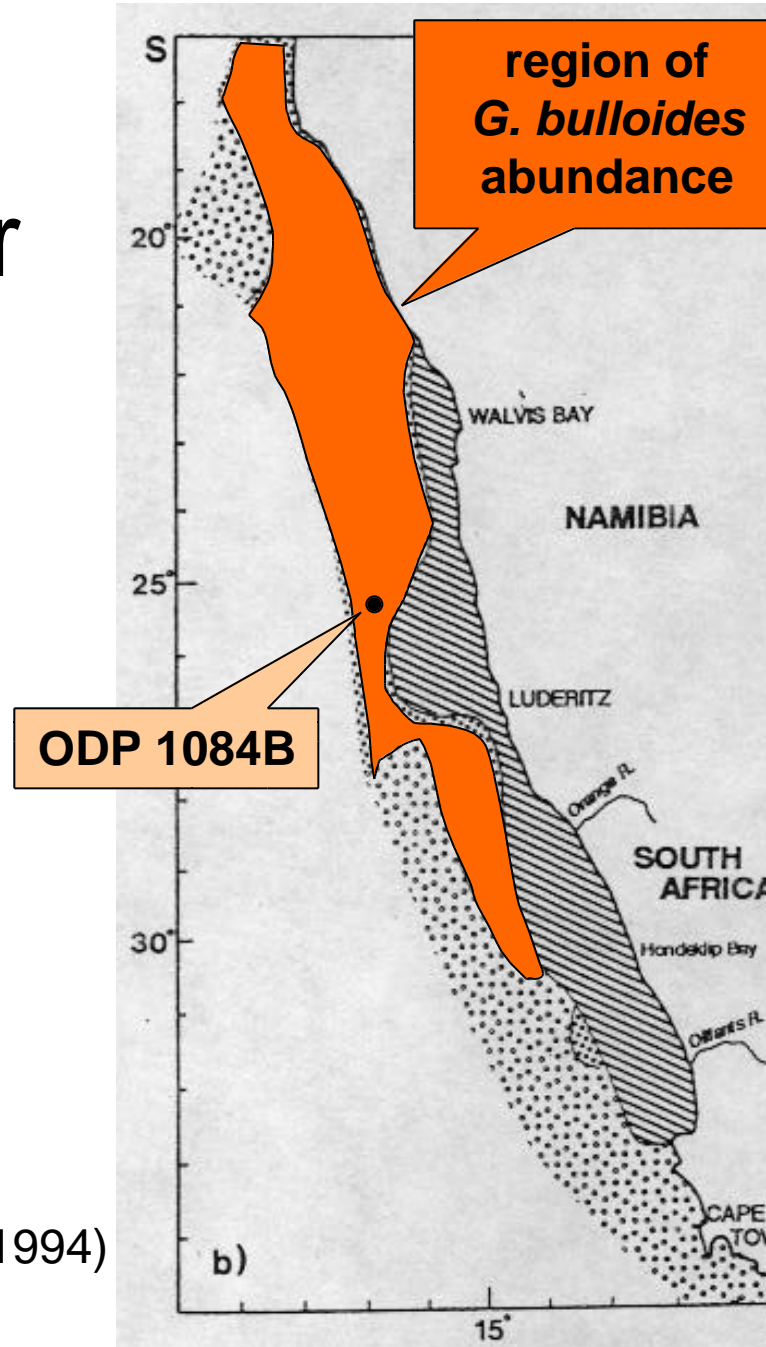
- protozoa that live in and above oceanic thermocline (upper ~400m)
- transitional to polar locations, upwelling environments
- $\text{CaCO}_3$  skeleton provides temperature, global ice volume, faunal abundance, and other proxies

# *G. bulloides*: upwelling indicator

Intermediate between  
species that prefer  
colder and warmer  
waters

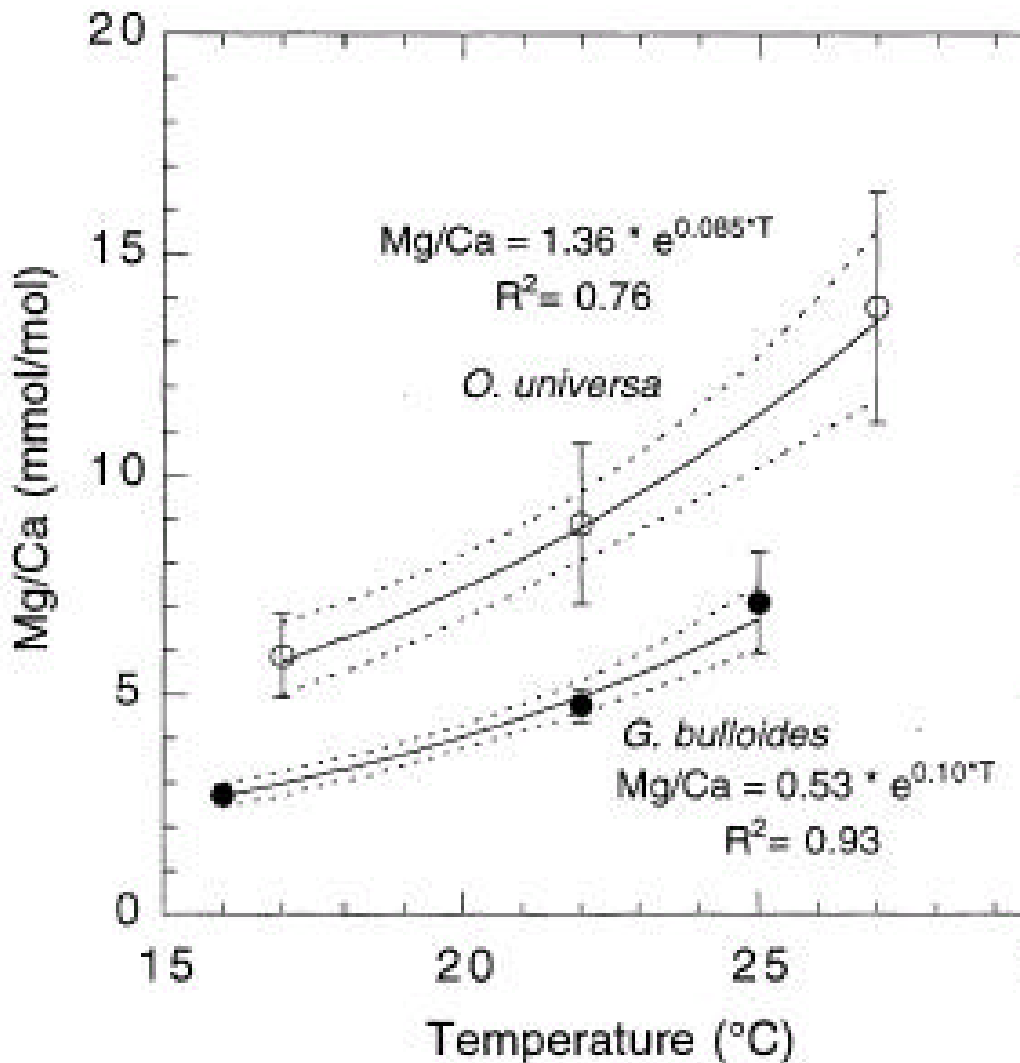
Faunal abundance  
data suggests  
movement of front?

(Giraudeau & Rogers, 1994)

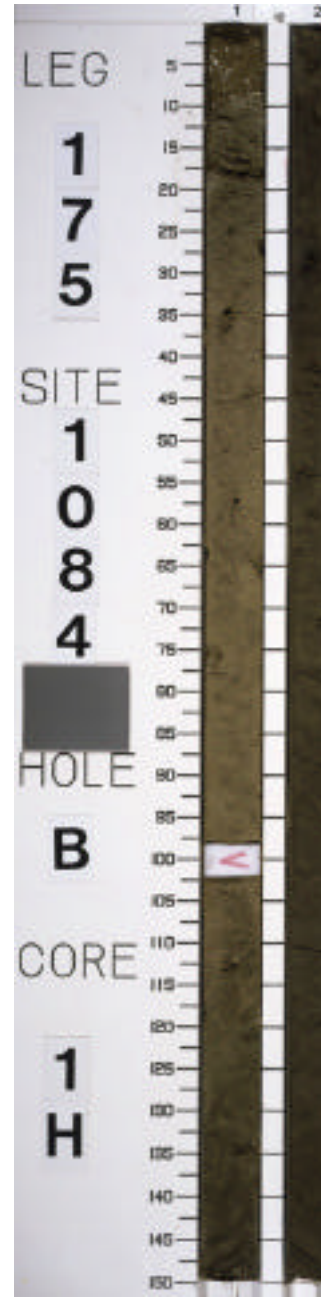
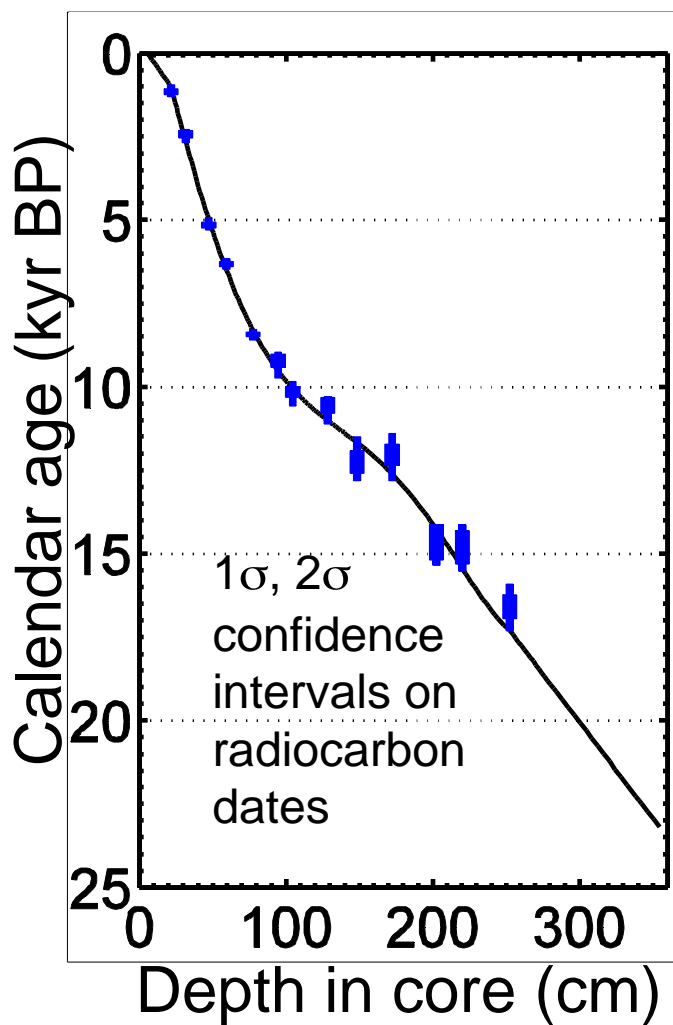


# ***G. bulloides*: temperature proxy**

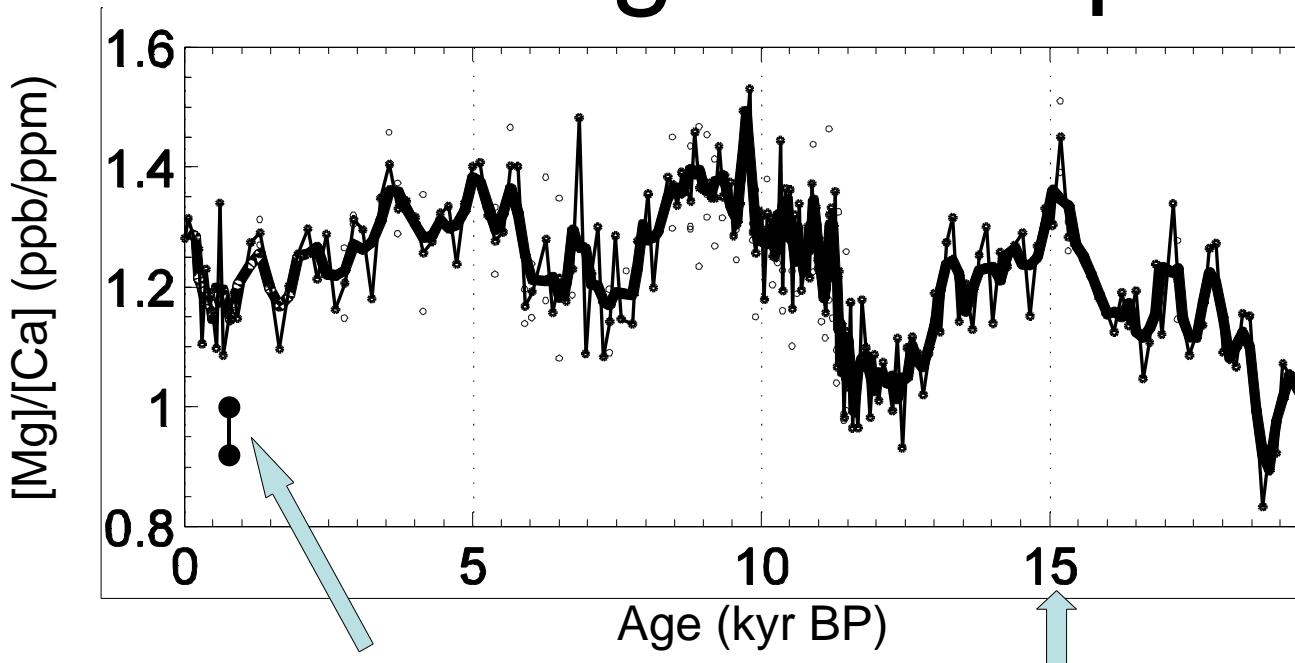
- Lea et al. 1999 growth experiments
- [Mg]/[Ca] of *G. bulloides* depends on temperature



# $^{14}\text{C}$ Age model for core 1084B:

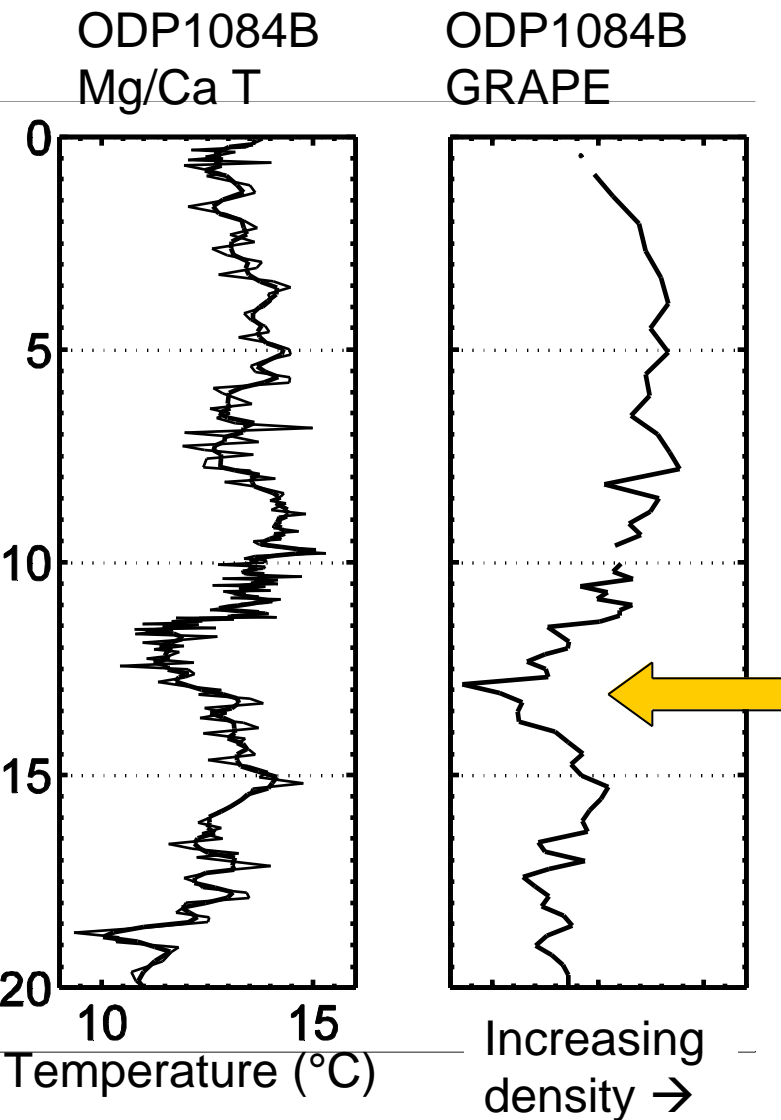


# Results: Mg/Ca temperature



- average difference between replicates: 0.08 (ppb/ppm)
- coretop value (1.3=13.6°C) matches modern winter (14.8°C) within confidence limits of Lea 1999 regression (+/-1.1°C)
- age model below ~15k yBP is not reliable: more dates needed
- wind strength fluctuation changes temperature, but how much is also due to changes in ocean currents? → further work

# MST GRAPE data:



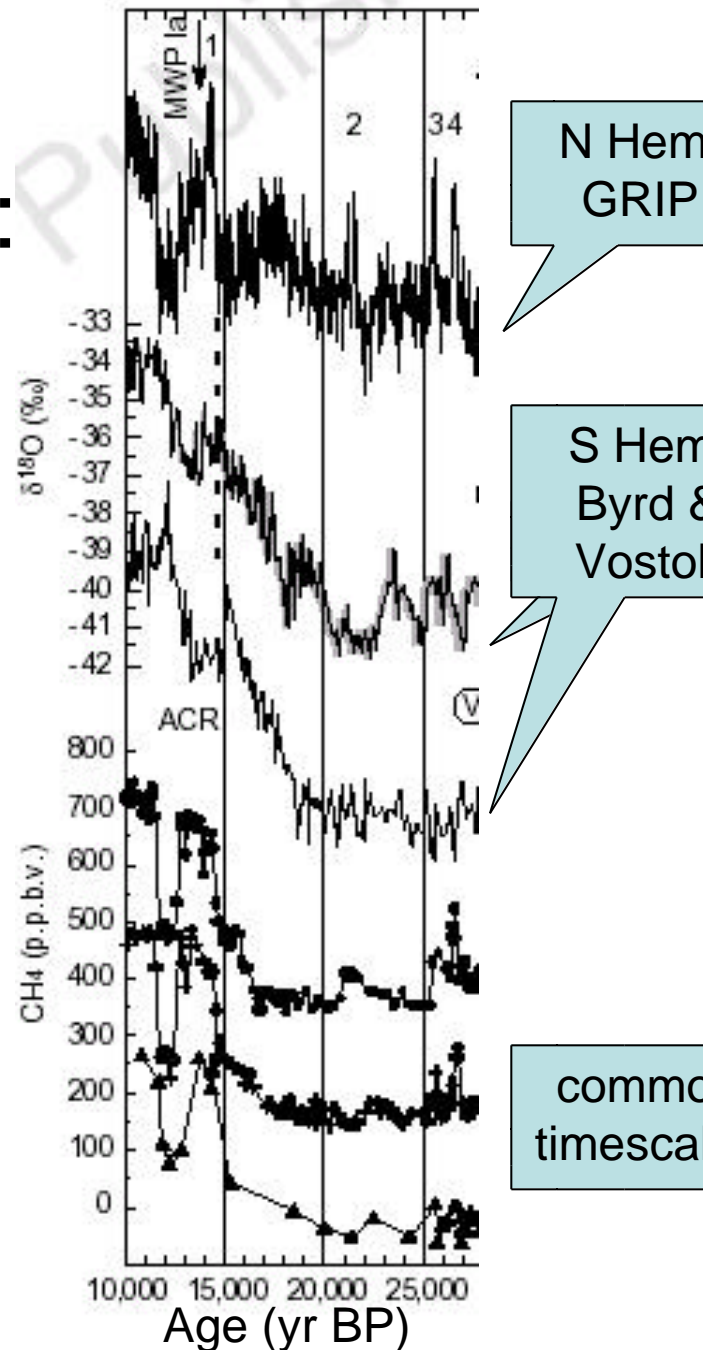
- Gamma Ray Attenuation inversely proportional to sediment density
- sharp density change at beginning of Younger Dryas (11,500-13,000 yr BP)



# Deglacial timing between hemispheres:

- N Hem: “Younger Dryas” cold period between glacial maximum and Holocene
- S Hem: “Antarctic Cold Reversal” precedes YD by ~1500 yrs, meltwater pulse by ~1000 yrs

(Blunier et al., 1998)



# Younger Dryas:

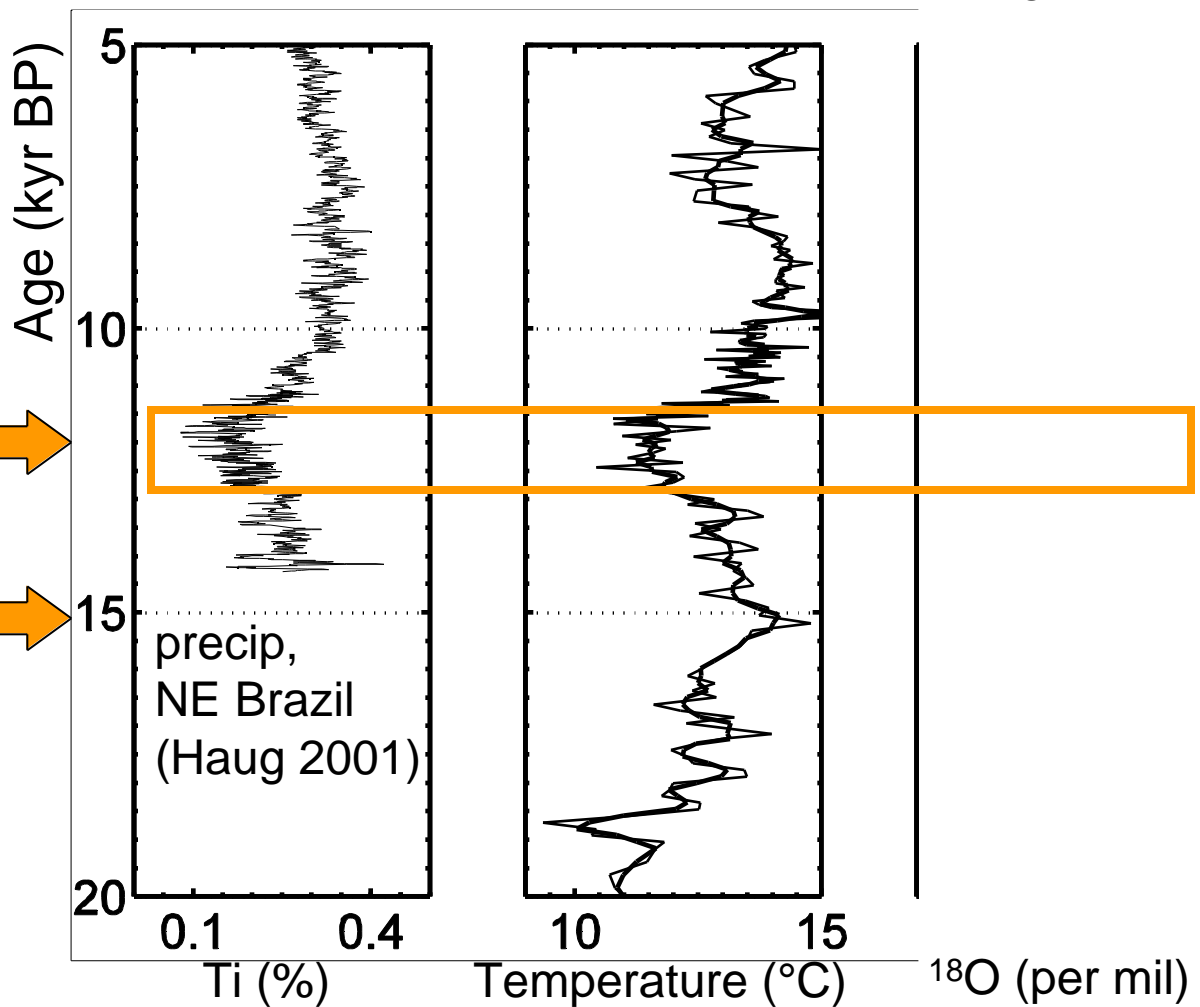
temperature,  
Benguela  
(ODP1084B)

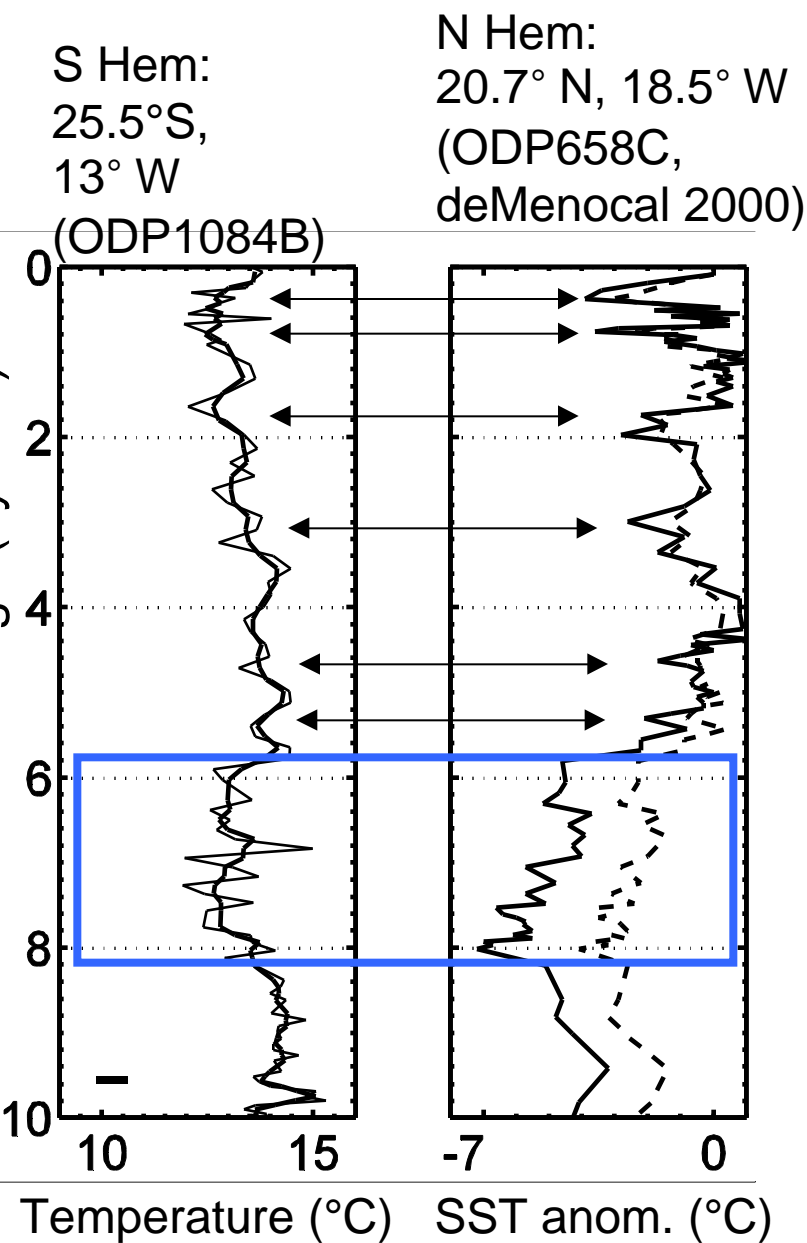
temperature,  
Greenland  
(Dansgaard 199

1084B  
matches  
N Hem  
records



S Hem  
ACR  
started  
15kyr BP





# Holocene:

- both hemispheres have same timing for subtropical Atlantic cooling events ~8kyrBP & 6.5kyrBP
- other millennial events seem to match too, but resolution too poor to say for sure

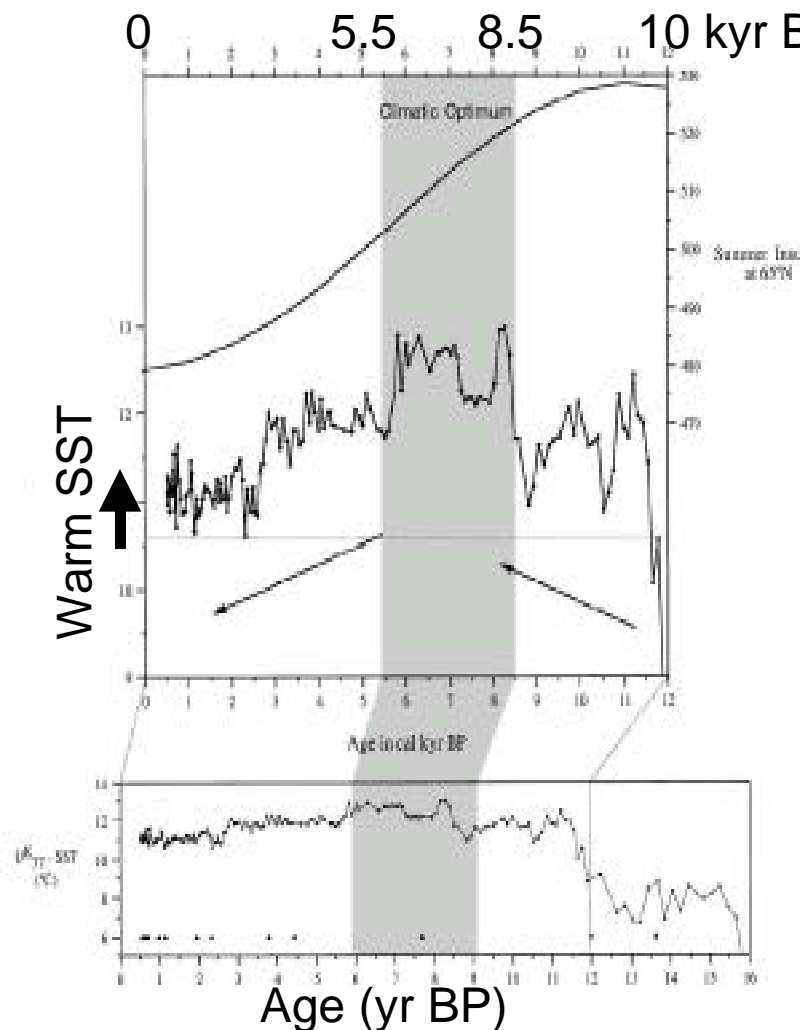


# mid-Holocene: 1084B cooling while N Atlantic warming

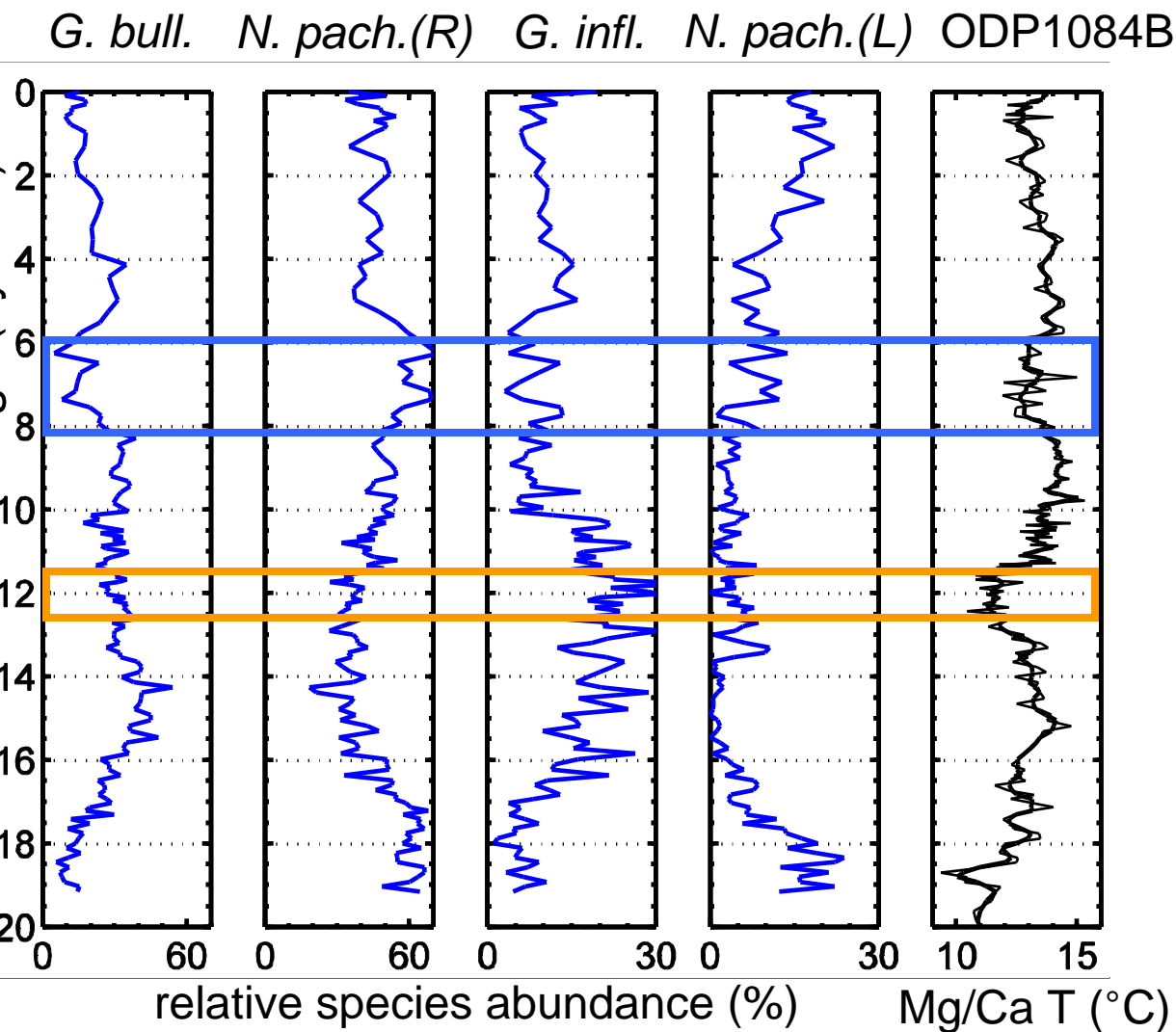
Norwegian Sea (65N) warms while Benguela upwelling region is cooling

Why is site 1084B in phase with subpolar N Hem for Younger Dryas but not 5.5-8.5kyr BP?

(Calvo et al., 2002)



# Faunal abundance results



- intriguing possibilities for faunal proxies...

# Faunal indicators:

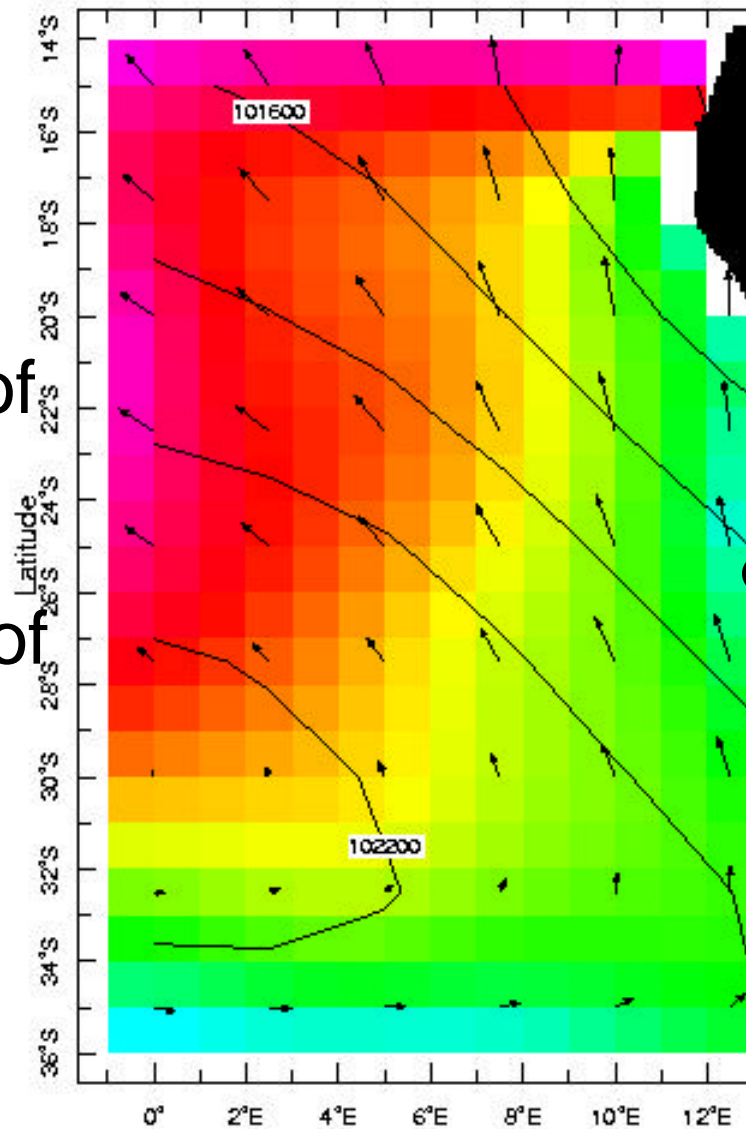
$[(L/B)]:$

W-E extension of upwelling cells?

$[(B+R)/(I+L)]:$

N-S movement of oceanic front?

Change in intensity of “thermohaline circulation”?



ODP 10841

NOAA WOA, NCEP-NCAR)

0.0 m/s

13

15

17

19



# Faunal indicators:

$[(L/B)]:$

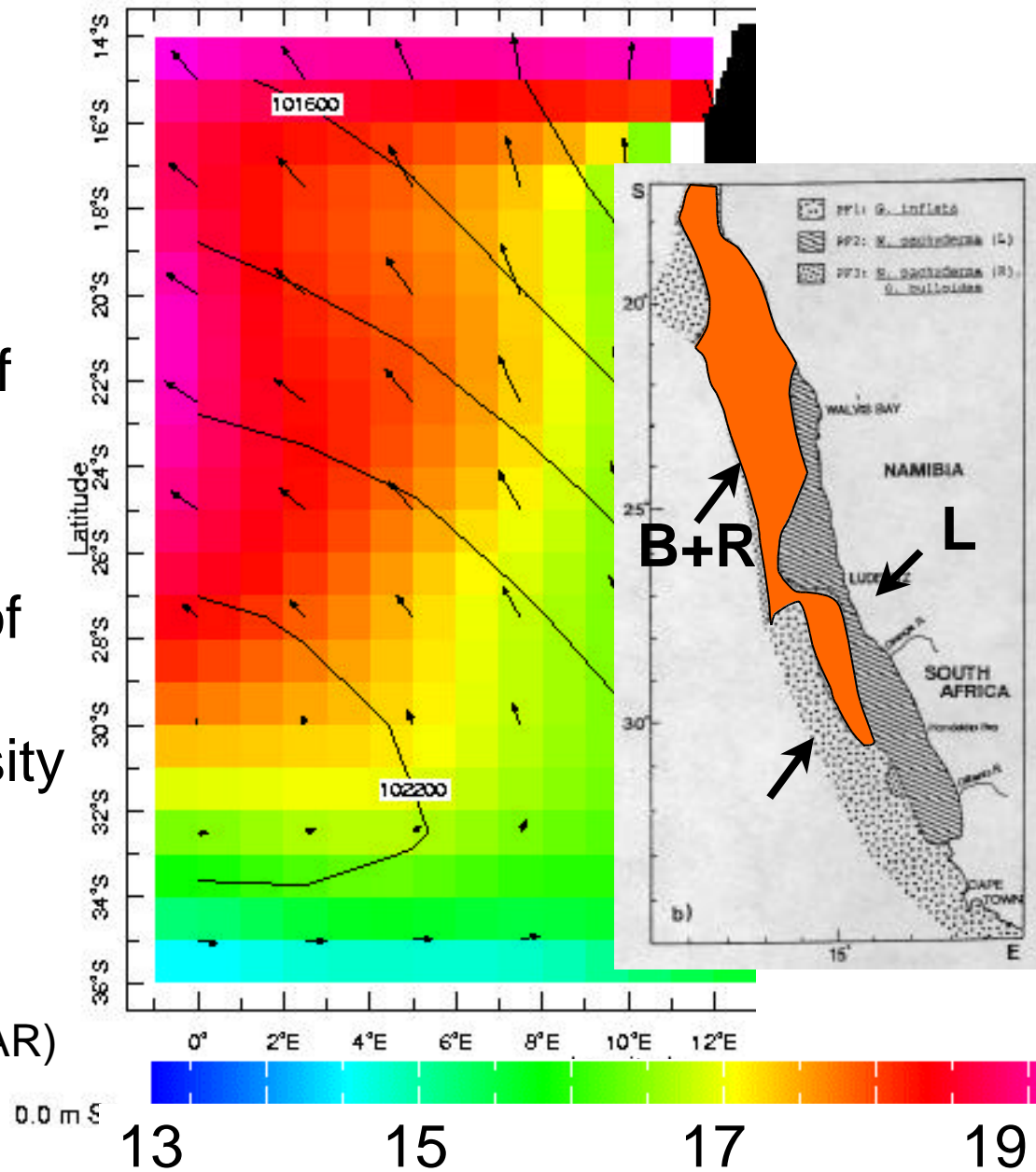
W-E extension of upwelling cells?

$[(B+R)/(I+L)]:$

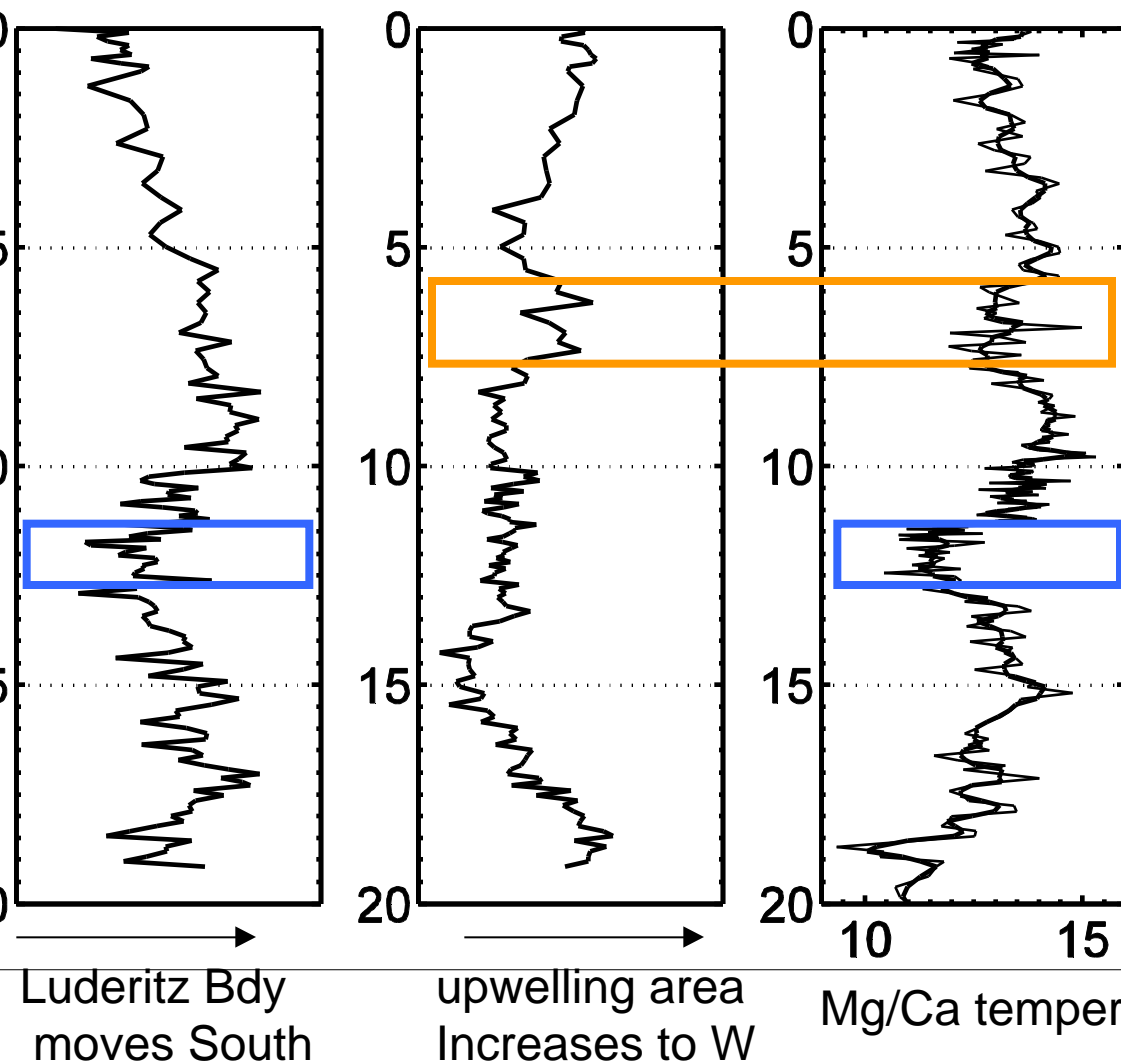
N-S movement of oceanic front?

Change in intensity of “thermohaline circulation”?

NOAA WOA, NCEP-NCAR)



# Why Hem phasing is different for two major climate events in ODP1084B:



- perhaps Holocene event is caused by changing wind strength  
→ Hems in phase
- while Younger Dryas is caused by changing ocean circulation  
→ Hems out of phase

# Conclusions:

## **S Hem subtropics:**

- match N Hem deglacial timing, not S Hem
- match N Hem subtropical mid-Holocene cooling events, but opposite of warming in subpolar N Atlantic

→ resolution in this core not quite high enough to show millennial variability

# Further Work:

- more  $^{14}\text{C}$  dates from lower section
  - $^{15}\text{N}$ , upwelling proxy
- future coring cruise is planned: hopefully it will get better sediments!!



*Thanks!*

